Jurassic-Pliocene biogeography: testing a model with velvet worm (Onychophora) vicariance

Julián Monge-Nájera

Centro de Investigación General, UNED, Costa Rica. Mailing address: Biología Tropical, Universidad de Costa Rica, San Jo sé, Costa Rica, telefax (506) 207 5550, julianm@cariari.ucr.ac.cr; jnajera@spiraxis.inbio.ac.cr.

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Abstract: Summary paleomaps of global continental vegetation from the Jurassic through the Pliocene were prepared (based on the literature) and used to define an area cladogram. Bouvier's natural classification of velvet worms (phylum Onychophora) was used to define a taxonomic/geographic cladogram of onychophorans for all regions where they are found today. Both cladograms show the same sequence of geographic vicariance. Thus, the paleogeographic model is supported by the taxonomic evidence. The paper includes a color atlas.

Key words: Onychophora, vicariance, retrovicariance, biogeography, cladistics, biogeography, Mesozoic, Cenozoic.

In the last decades, the availability of computers has resulted in a rapidly growing number of paleogeographic models that often differ among themselves (Scotese *et al.* 1988, Gyllenhaal *et al.* 1991, Patzkowsky *et al.* 1991) and normally use little or no biological information as a basis (Briggs 1989, Parrish 1993a).

This paper, which emphasizes the tropics and is addressed mainly to biogeographers, presents a summary of several paleogeographic reconstructions and includes brief descriptions of terrestrial ecosystems from a recent and comprehensive review (Behrensmeyer *et al.* 1992). The maps were used to produce an area cladogram that shows the sequence in which large landmasses separated from the Jurassic through the Pliocene, for comparison with independent evidence based on velvet worm taxonomy (phylum Onychophora).

Onychophorans, a group of scarce terrestrial worms, have long been considered appropriate biogeographic subjects because of their old age and low capacity to disperse (Sedgwick 1908, Vachon 1954, Monge-Nájera 1995). A review of the classification and keys of the phylum presented by Bouvier in 1905, showed that he intended them to present a natural classification (see Bouvier 1905: 7, 74-66). Thus, for this study, his classification was converted to graphic form by preparing a cladogram, which in this group is thought to reflect both the sequence of geographic isolation and of speciation (see Sedgwick 1905 and Ruhberg 1985).

This paper shows that when both cladograms are compared, they imply the same sequence of geographic isolation for the general areas where the phylum occurs today.

MATERIAL AND METHODS

Paleogeographic reconstructions are based on the projections presented by Briggs (1989), who takes into account biological evidence, and on a list of sources too long to repeat here but presented in Monge-Nájera (1995). Marine currents (Schopf 1980) and climatic data (Condie 80, Parrish 1993b) were added to the original maps. The hypotheses about overall vegetation types are based on the general associations between temperature, precipitation and plants in the modern world (Walter and Breckle 1986).

To convert the maps into the area cladogram, the sequence of separations was represented by a line that bifurcated whenever two areas separated. The same procedure was used for the onychophoran cladogram, by bifurcating the line whenever cladogenesis took place according to the classification and identification keys presented by Bouvier (1905). This is the standard procedure to make classifications into cladograms and assumes them to be natural rather than artificial (Brusca and Brusca 1990).

RESULTS

The paleogeographic maps appear in Figs. 1-11 and the area cladograms in Fig. 12.

Area relationships: There is important disagreement among authors regarding the position of India in the late Jurassic; it is thought to be closer to Australia, isolated, or connected with Antarctica according to fossil evidence from belemnites, dinosaurs and plants, respectively (Hallam 1973, 1994, Schuster 1983, Briggs 1989). The position presented in Fig. 3 and the following maps is the one that seems less conflictive with current evidence.

Despite the separation shown between Europe and Asia in the early Cretaceous (Fig. 4), plants and dinosaurs were relatively cosmopolitan in Laurasia and in Gondwana (Schuster 1976, Behrensmeyer *et al.* 1992). Similarly, North America may have actually been separated from Europe, in contrast with this reconstruction (Burden and Langille 1991). A possible connection of the Greater Antilles with Yucatan or Florida is debated (Coney 1982, Iturralde-Vinent 1988), and India was more to the south according to Barron (1987).

For the mid Cretaceous, reconstructions differ significantly about the position of India and the relationship between South America and Africa (Barron 1987, Dott and Batten 1988). In contrast with this map (Fig. 5), some authors believe that in the late Cretaceous Greenland was becoming an island and that the Antilles were more to the west, in the area where the early Central America is shown here (Alvarado 1988, Dott and Batten 1988). Again, there are other opinions about the position of India (Briggs 1989).

In the Paleocene, this map (Fig. 7) differs from those of Dott and Batten (1988), and Piccoli *et al.* (1991) because they place India to the south and Australia more separated from Antarctica. Also in disagreement with the map of Fig. 8, India may have been still an island in the Eocene (Scotese *et al.* 1988).

In the Oligocene (Fig. 9), South America could have been closer to Africa (Alvarado 1988), but for the more recent periods (Figs. 10 and 11) there are no significant disagreements between the consulted reconstructions.

Vegetation: This Jurassic reconstruction is consistent with Mojon and Mouchet (1992) and with Colorado data in Parrish (1993), but not with paleontological data which indicate moister conditions in NE North America (McDonald 1992), France (Thevenard 1993) and Denmark (Arndorff 1993).

There is a better fit for the Cretaceous (Antartic Peninsula, Chapman and Smellie 1992; Belgium, Bless et al. 1993; Nigeria, Edet and Nyong 1993; the Andes, Gayet et al. 1993 and Tibet, Xu and Mao 1992). Nevertheless, conifers grew in many parts of Laurasia that this map depicts as shrubland (Cevallos-Ferriz 1992, Millar 1993). Although no recent fossil data for the Paleocene were found, for the Eocene the fit was also satisfactory (Cuba, Sieglne-Farkas et al. 1992; France, Kohler 1992; Argentina, Leguizamon et al. 1993) with the exception of parts of Saudi Arabia and France that had mangroves and humid forests, respectively, instead of savannas (Srivastava and Binoa 1991).

For the Oligocene there is full correspondence between the map and fossils for sites so separate as Antarctica and Europe (Alvinerie *et al.* 1992, Barron and Mahood 1993, Knobloch 1993).

A greater number of recent studies refer to the Miocene and are consistent with the map for Chile (Troncoso 1991), Hungary (Hably 1992), India (Prasad 1993) and Japan (Ina 1992, Sato 1992, Yamanoi 1993). The map



Fig. 1. Identification of continents in Pangaea. Use as basis to identify landmasses in following figures.

shows xeric vegetation in two places where moist-type fossil vegetation was found: Turkey (Gemici *et al.* 1993) and the Niger Delta (Morley and Richards 1993).

The Pliocene reconstruction is fully consistent with data from Antarctica (Marchant *et al.* 1993), North America (Watkins 1992) and Europe (Morley and Richards 1993).

Area cladograms: The area cladogram based on paleogeographic reconstructions (Fig. 12 A) shows a geologic time scale for the general fragmentation process, including first the latitudinal separation of two large landmasses, followed by separation of Indomalaysia from the Neotropica-Equatorial Africa complex. Australasia separated from the Chile-South Africa complex.

The second cladogram (Fig. 12 B) is fully independent of the first because it is based on taxonomic evidence. It shows how the onychophorans became geographically and genetically isolated. Genetical isolation resulted in an speciation process (*i.e.* cladogenesis), according to the taxonomic work done early in the century by Bouvier (1905). This cladogram shows first a large division into the equatorial and southern families (respectively: Peripatidae and Peripatopsidae). Bouvier's (1905) classification has no quantification of the taxonomic **REVISTA DE BIOLOGIA TROPICAL**



Fig. 2. Early Jurassic (180-165 million years before Present, mybP). Broken line: limit of submerged coastal areas (several authors show more or less flooded area than these reconstructions; the same applies to the following maps). Key to climates: 1 very dry, 2 dry, 3 moist and 4 very moist. The tropical-subtropical belt is marked by latitudinal lines. Key to hypothetical dominant vegetation types: **Tropical regions:** 1 Desert and semidesert, 2 Steppes and shrubby vegetation, 3 Savanna and 4 Rain forest. **Temperate regions:** 1 and 2 similar to tropical regions, 3 temperate forest and grassland, and 4 Coniferous forest and other types of temperate forest. Near the poles, only the drier areas had tundra and alpine vegetation.

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Fig. 3. Late Jurassic (160-135 mybP). Although not specifically showed in this and following maps, an important current is thought to have moved towards the west in the area between South and North America. Key to climate and vegetation as in Fig. 2. Summary of Jurassic ecology: **Climate and Flora:** Warm arid and seasonally arid areas, particularly near Tropic in W. Pangaea; moister in China and other high latitudes; breakdown of monsoonal system. No impassable barriers between continents, only minor N.-S. provincialism. At least in early Jurassic, angyosperms rare. Four defined floral provinces: N. Laurasia, S. Laurasia, N. Gondwana, S. Gondwana. Tropics: seasonally dry. Angyosperms more important near tropics. **Fauna:** Coleoptera very varied; first Thysanoptera. Large, high-browsing sauropods dominant; probably plants with more defenses and large re-growth capacity despite aridity. Some herbivores fed 10-12 m above ground.



Fig. 4. Early Cretaceous (135-110 mybP). Key to climate and vegetation as in Fig. 1.

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Fig. 5. Mid Cretaceous (110-90 mybP). Key to climate and vegation as in Fig. 1.



Fig. 6. Late Cretaceous (89-65 mybP). Key to climate and vegetation as in Fig. 2. Summary of Cretaceous ecology. Climate and flora: Flowering plants appear with high diversity (specially riparian and burnt habitat) but were not important part of forest biomass. K/T boundary (Cretaceous / Tertiary): great extinction of large terrestrial vertebrates associated with plant devastation; less extinction apparent toward S. Pole. Fauna: Early period: Micropterigid lepidopterans; larvae fed on moss, liverworts and detritus: adults maybe ate polen. Late period: Lepidopteran miners, Orthoptera (Acridoidea); Hymenoptera (Cynipidae), aphids, Isoptera, ants. Laurasia: abundant high-browsing sauropods. Gondwana: sauropods dominant; mammals rare. Abundant small mammalian herbivores in Campanian-Mastrichtian (83-65). S. American mammals highly endemic. Hadrosaurid ornithopods and ceratopsicians important, oral food processing. Herd foraging? Coevolution with plants? Most feeding 1-3 m above ground.

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Fig. 7. Paleocene (65-54 mybP). Key to climate and vegetation as in Fig. 2. Summary of ecology. Climate and Flora: Colder climate? Probably angiosperms suffered great decimation (recovered near end of Paleocene) and were substituted by fern-dominated community. Deciduous conifer forest, but new Argentinan fossils indicate tropical and montane rain forest and savanna-sclerophyllous forest. Fauna: Rise of mammals, many consumers of fruit, seed and small prey; specialized browsers not important. Coevolution of plants and their mammal dispersers. Shift from browsers to fruits and seeds.



Fig. 8. Eocene (54-38 mybP). Key to climate and vegetation as in Fig. 2. Summary of ecology. **Climate and Flora:** Drier than Paleocene.; 33 mybP cooler, more seasonal toward end. Steep latitudinal thermal gradients. Final upwelling of Rocky Mountain rain-shadow: "Terminal Eocene event". Golden age of tropicality and angiosperm-dominated closed forest. Rainforest up to 30⁰N, subtropical to 60⁰N. Complexity of communities equal to highest today. Menispermaceae & Icacinaceae lianas. Australia: angiosperm dominated tropical and subtropical rainforest, *Eucalyptus*. Tasmania: cooler, subtropical, rainy. Argentina: moist with seasonal dryness. **Fauna:** Modern moths (Monotrysia) and butterflies (Ditrysia). Formicoid ant radiation. Contact of European and N. American mammals resulted in: (1) turnover, (2) homogenization; trend to larger size. Toward end: more open habitat fauna (more lophodonty and cursoriality; less arboreal forms), especially away from tropics. In Oligocene-Miocene (23 myb^P) partial return to browzing.



Fig. 9. Oligocene (38-23 mybP). Key to climate and vegetation as in Fig. 2. Summary of ecology: Climate and flora: Cooler, more seasonal (result of new ice caps?), drier. Cooling began earlier in somer areas. Monsoonal rainfall in Africa. Microthermal, broad leaved forest over large regions of Northern Hemisphere. Later: Grasses and composites increase in abundance. North America: stronger rain shadow caused by Rocky Mts., vulcanism in Pacific N.W. In S. America: humid forests common; N. Africa: seasonally dry forest; Australia: plants adapted to sandy and water-stressed soils. Acacias & scherophyllous forests grow. Fauna: Large-bodied perissodactyls associated with felid-, canid- and hyena- like carnivores that were slower than modern forms. Africa: diverse micromammals, large proboscideans, hyracoids fill mid-size browser niche; fauna of Asia connected with Europe. Australia: marsupials greatly expand.



Fig. 10. Miocene (23-5 mybP). Key to climate and vegetation as in Fig. 2. Summary of ecology: Climate and flora: Circum-Antartic marine current well established. Mid-Miocene global warming; Andes rise, causing rain-shadow; gradual cooling and drying in parts of America. Cool-warm cycles in S. America. 22 mybP: Africa, Arabia and Eurasia connected by intermitent bridges, less moisture from Thethys, more rain-shadow in E. Africa, more seasonal aridity (13 mybP); oscillations of cool and warm climate(9 mybP), drier Mediterranean. 18 mybP: Broad-leaved evergreen vegetation and coniferous forests expanded in northern hemisphere. Broad-leaved deciduous forest reduced. 13 mybP: Savannas and grasslands grow, forest replaced by woodlands in some areas. S. America: mixed forest and savanna as in Africa. Australia: wind & insect pollinated plants increase, rain forest replaced by dry forest and woodland. Fauna: Aprox. 12 mybP: diversification of small mammals (especially rodents), open-vegetation herbivores, large mammalian carnivores, and snakes. Second half of Miocene: diversification of large, soaring, carnivorous birds and other open-habitat birds, higher-crowned teeth in horses; increased cursoriality and body size, large head size; grazing, artiodactyls diversify. Biotic immigration Eurasia- N. America and exchange Eurasia-Africa. Australia: rain forest support large, slow moving browsers, flightless birds, lion and wolf-like carnivores.



Fig. 11. Pliocene (5-2 mybP). Key to climate and vegetation as in Fig. 2. Summary of ecology: Climate and flora: Himalayan uplift. Eurasia: hotter weather. 2.4 mybP: Periodic glaciation and build up of Northern Hemisphere ice sheets. Isthmus of Panama emerges. Retreat of paleotropical flora in N. hemisphere; expansion of temperate deciduous trees, grasses, composites. Australia drier, more wet-arid fluctuations. Conifers increase at high latitudes. Prairies, steppes and grasslands in N. America. Increased diversity (especially of rodents). Fauna: Further diversification of large, cursorial, grazing herbivores associated with evolution of larger, faster carnivores. Diversification of cryptic rodents, birds and small carnivores. Bipedal primates. Major invasion of S. America by N. American mammals; extinction of marsupial carnivores; Eurasia: more savanna fauna. Africa: more small carnivores, first large felids. Australia: open habitats inhabited by kangaroos; burrowing wombats.



Fig. 12. Area cladograms produced from the reconstructions in Figs. 2-11 (A, scale bar: 50 million years) and from the systematics of velvet worms (B) according to Bouvier (1905). Both cladograms show the same sequence of geographic isolation. Cladogram characters: 1. Sexual orifice between penultimate leg pair, brown reddish pigment soluble in preservation fluid, 14 regular dorsal folds per segment (rarely 22), variable number of foot-soles which are of subequal length, urinary pore of fourth and fifth leg pairs under the middle rear part of the third sole arch (or under it), inner mandible blade with a saw and a diastema; salivary reservoirs, seminal receptacles and ovular receptacles present. 2. Sexual orifice amid the last leg pair or behind it, pigmentation often greenish blue not greatly altered by preservation fluids, very irregular dorsal folds of which there are 14-25 per segment, three-arched foot soles (middle ones much larger), urinary pore of fourth and fifth leg pairs amid the third arch, inner mandible blade without saw and without diastema, salivary reservoirs absent. 3. At least three foot papillae, pores of male annal glands ventraly isolated in front of the anus, very small endogenous eggs, placented embryos. 4. Very irregular dorsal folds, rudimentary seminal receptacle, eggs with very reduced reserves.

distance between groups, and for that reason the cladogram has the same distances between branch origins. Nevertheless, it does show that the species of Indomalaysia were the first to become isolated in the equatorial family and that the species of Neotropica separated more recently from those of Equatorial Africa. In the south, Australasian species were the first to separate and those of Chile were closely related with South African species until more recent times.

DISCUSSION

Although paleontological remains may provide more detailed paleoecological information than geophysical evidence, they are subject to important errors in identification, interpretation (Crame 1992, Varol 1992, Head 1993, Tanai 1993) and dating (Burnett et al. 1992, Rees 1993). The basic divergence in a comparison of the vegetation maps presented here (based on temperature and rainfall) with paleontological data is that fossils from moist habitats have been found in some areas where the maps indicate savannas or similar biomes. If the fossil data are more reliable than the climatic models. which seems probable, this may reflect the ocurrence of highlands not considered in this reconstruction, as well as the effects of lakes, rivers and seashores. Future maps could improve on this condition by including more topographic details. This paper emphasized differences because there is more disagreement among authors of paleogeographic reconstructions than is normally realized. Detailed local studies (e.g. Seyfried and Hellmann 1994) are also unveiling valuable new information.

The fit of area cladograms produced independently from the systematics of velvet worms and from tectonic analysis corrected with paleontological evidence, supports this reconstruction on a very general level. However, this fit should not be taken as support for every detail of the maps, which are only a rough representation of what must have been a much more complex paleogeographic history.

A previous attempt of relating paleomaps with onychophoran distribution (Monge-Nájera 1995), which lacked taxonomic and vegetation information, produced less clear results for the Indomalaysia-Australasia area and showed a different isolation sequence within the family Peripatidae. The paleogeographic reconstruction presented here is consistent with the taxonomists' association of South African and Chilean species (Bouvier 1905, Ruhberg 1985), highlighting the importance of including vegetation in this type of study. A further advantage of this taxonomic-paleogeographic model is that it predicts more specific geologic times for the genetic isolation of onychophorans inhabiting different parts of the world. Such predictions are testable with the use of biochemical techniques that measure genetic distance (see Walker 1995). I hope that such studies will be undertaken by biochemical systematists in the future.

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RESUMEN

Con base en la literatura se hizo mapas resumidos de la vegetación continental desde el Jurásico hasta el Plioceno y se usaron para definir un cladograma de áreas. Independientemente, se usó la clasificación natural de los gusanos del filo Onychophora (elaborada por Bouvier) para preparar un cladograma taxonómico y geográfico de todas las regiones ocupadas por ellos en la actualidad. Ambos cladogramas muestran la misma secuencia de separación geográfica, por lo que el modelo paleogeográfico es apoyado por la evidencia taxonómica. El artículo incluye un atlas en color.

REFERENCES

Alvarado, G.E. 1988. Centroamérica y las Antillas: Puente, barrera y filtro biológico entre Norte y Sudamérica (Cretácico al Presente). Geoistmo (Costa Rica) 2: 9-25.

- Alvinerie, J., M.T. Antunes, B. Cahuzac, A. Lauriatrage, C. Montenat & C. Pujol. 1992. Synthetic data on the paleogeographic history of Northeastern Atlantic and Betic-Rifian basin, during the Neogene (from Brittany, France, to Morocco). Palaeogeogr. Palaeoclimat. Palaeoecol. 95: 263-286.
- Arndorff, L. 1993. Lateral relations of deltaic palaeosols from the Jurassic Ronne Formation on the island of Bornholm, Denmark. Palaeogeogr. Palaeoclimatol. Palaeoecol. 103: 235-250.
- Barron, E.J. 1987. Global Cretaceous paleogeography. Palaeogeogr. Palaeoclimat. Palaeoecol. 59: 207-214.
- Barron, J.A. & A.D. Mahood. 1993. Exceptionally wellpreserved early Oligocene diatoms from glatial sediments of Prydz Bay, East Antartica. Micropaleontology 39: 29-45.
- Bless, M.J., M.M. Dusar, P.J. Felder & R. Swennen. 1992-1993. Lithology and biostratigraphy of Upper Cretaceous-Paleocene carbonates in the Molenbeersel borehole (NE Belgium). Geol. Mijnbouw 71: 239-257.
- Behrensmeyer, A.K., J.D. Damuth, W.A. DiMichele, R. Potts, H.-D. Sues & S.L. Wing. 1992. Terrestrial Ecosystems Through Time. University of Chicago, Illinois. 568 p.
- Briggs, J.C. 1989. The historic biogeography of India: Isolation or contact? Syst. Zool. 38: 322-332.
- Brusca, R.C. & G.J. Brusca. 1990. Invertebrates. Sinauer, Sunderland, Massachusetts. 922 p.
- Burden, E.T. & A.B. Langille. 1991. Palynology of Cretaceous and Tertiary strata, northeast Baffin island, Northwest Territories, Canada: Implications for the history of rifting in Baffin Bay. Palynology 15: 91-114.
- Burnett, J.A., J.M. Hancock, W.J. Kennedy & A.R. Lord. 1992. Macrofossil, planktonic foraminiferal and nannofossil zonation at the Campanian/Maastrichtian boundary. Newsl. Stratigr. 27: 157-172.
- Carpenter, R. & A.M. Buchanan. 1993. Oligocene leaves, fruit and flowers of the Cunoniaceae from Cethana, Tasmania. Aust. Syst. Bot. 6: 91-109.
- Cevallos-Ferriz, S.R.S. 1992. Tres maderas de gimnospermas del Cretácico del norte de México. An. Inst. Biol. Univ. Nac. Autón. Mexico (Ser. Bot.) 63: 111-137.
- Chapman, J.L. & J.L. Smellie. 1992. Cretaceous fossil wood and palynomorphs from Williams Point, Livingston Island, Antarctic Peninsula. Rev. Palaeobot. Palynol. 74: 163-192.
- Condie, K.C. 1982. Plate tectonics and crustal evolution. Pergamon, New York. 310 p.

- Coney, P.J. 1982. Plate tectonic constraints on the biogeography of Middle America and the Caribbean region. Ann. Missouri Bot. Gard. 69: 432-443.
- Crame, J.A. 1992. Late Cretaceous palaeoenvironments and biotas: An Antarctic perspective. Antarct. Sci. 4: 371-382.
- Deschamps, C.M. & E.P. Tonni. 1992. Vertebrados del Pleistoceno Tardío-Holoceno de Arroyo Naposta Grande, Provincia de Buenos Aires: características paleoambientales. Ameghiniana 29: 201-210.
- Dmitrenko, O.B. 1992. The biostratigraphic calcareous nannoplankton association from the bottom sediments of the eastern Pacific subequatorial area [in Russian]. Okeanologiya 32: 761-768.
- Dott, R.H., Jr. & R.L. Batten. 1988. Evolution of the Earth. McGraw-Hill, New York. 120 p.
- Edet, J.J. & E.E. Nyong. 1993. Depositional environments, sea-level history and palaeobiogeography of the late Campanian-Mastrichtian on the Calabar flank, SE Nigeria. Palaeogeogr. Palaeoclimatol. Palaeoecol. 102: 161-175
- Gemici, Y., E. Akyol & F. Akgun. 1993. Macro and microfossil flora of the Sahinali (Aydin) Neogene basin. Turk. J. Bot. 17: 91-106.
- Gruas-Cavagnetto, C. & E. Kohler. 1992. Fossil polen of Euphorbiaceae from the Eocene of France. Grana 31: 291-304.
- Gyllenhaal, E.D., C.J. Engberts, P.J. Markwick, L.H. Smith & M.E. Patzkowsky. 1991. The Fujita-Ziegler model: a new semiquantitative technique for estimating paleoclimate from paleogeographic maps. Palaeogeogr. Palaeoclimat. Palaeoecol. 86: 41-66.
- Hably, L. 1992. Early and late Miocene floras from the Iharosbereny-I and Tiszapalkonya-I boreholes. Fragm. Mineral Palaeontol. 15: 7-40.
- Hallam, A. (ed.). 1973. Atlas of Palaeobiogeography. Elsevier, Amsterdam. 531 p.
- Hallam, A. 1994. An outline of Phanerozoic biogeography. Oxford University, Oxford. 200 p.
- Head, M.J. 1993. Dinoflagellates, sporomorphs, and other palynomorphs from the Upper Pliocene St. Erth beds of Cornwall, southwestern England. Paleontol. Soc. Mem. 31: 1-62.
- Horton, A., D.H. Keen, M.H. Field, J.E. Robinson, G.R. Coope, A.P. Currant, D.K. Graham, C.P. Green & L.M. Phillips. 1992. The Hoxnian Interglacial deposits at Woodston, Peterborough. Phil. Trans. Roy. Soc. London (B: Biol. Sci.) 338: 131-164.
- Ina, H. 1992. Miocene vegetational and climatic history of the eastern part of the Setouchi Geologic Province, Japan. J. Earth Sci. Nagoya Univ. 39: 47-82.

- Iturralde-Vinent, M.A. 1988. Naturaleza geológica de Cuba. Científico-Técnica, Havana. 146 p.
- Knobloch, E., Z. Kvacek, C. Buzek, D.H. Mai & D.J. Batten. 1993. Evolutionary significance of floristic changes in the Northern Hemisphere during the Late Cretaceous and Palaeogene, with particular reference to Central Europe. Palaeobot. Palynol. 78: 41-54.
- Leguizamon, R., C.G. Asato & H.H. Camacho. 1993. Calcareous algae from arroyo verde formation (Eocene, Chubut province, Argentina): Systematic description of the main genera and palaeoenvironmental considerations. Ameghiniana 30: 9-15.
- Lewis, D.W. 1992. Anatomy of an uncorformity on mid-Oligocene Amuri Limestone, Carterbury, New Zealand. N. Zeal. J. Geol. Geophys. 35: 463-475.
- McDonald, N.G. 1992. Paleontology of the Early Mesozoic (Newark supergroup) rocks of the Conneticut Valley. Northeast. Geol. 14: 185-199.
- Marchant, D.R., C.C. Swisher, D.R. Lux, D.P. West & G.H. Denton. 1993. Pliocene paleoclimate and East Antartic ice-sheet history from surficial ash deposits. Science 260: 667-670.
- Millar, C.I. 1993. Impact of the Eocene on the evolution of *Pinus* L. Ann. Mo. Bot. Gard. 80: 471-498.
- Mojon, P.-O. & P. Mouchet. 1992. Revision and validation of *Echinohara pecki*, new combination (Maedler) Grambast 1965 (Claveatoraceae, Charophyta) according to unpublished material in the lower Kimmeridgian in the northern Swiss Jura. Bull. Soc. Neuchatel Sci. Nat. 115: 9-21.
- Monge-Nájera, J. 1995. Phylogeny, biogeography and reproductive trends in the Onychophora. Zool. J. Linn. Soc. [London] 114:21-60.
- Parrish, J.T. 1993a. A brief discussion of the history, strengths and limitations of conceptual climate models for the pre-Quaternary time. Phil. Trans. R. Soc. Lond. B 341: 263-266.
- Parrish, J.T. 1993b. Climate of the supercontinent Pangea. J. Geol. 101: 215-233.
- Parrish, J.T. & R.L. Curtis. 1982. Atmospheric circulation, upwelling, and organic-rich rocks in the Mesozoic and Cenozoic eras. Palaeogeogr. Palaeoclimat. Palaeoecol. 40: 31-66.
- Parrish, J.T., A.M. Ziegler & C.R. Scotese. 1991. Rainfall patterns and the distribution of coals and evaporites in the Mesozoic and Cenozoic. Palaeogeogr. Palaeoclimat. Palaeoecol. 40: 67-81.
- Patzkowsky, M.E., L.H. Smith, P.J. Markwick, D.J. Engberts & E.D. Gyllenhaal. 1991. Application of the Fujita-Ziegler paleoclimate model: Early Permian and Late Cretaceous examples. Palaeogeogr. Palaeoclimat. Palaeoecol. 86: 67-85.

- Piccoli, G., S. Sartori, A. Franchino, R. Pedron, L. Claudio & A.R. Natale. 1991. Mathematical model of faunal spreading in benthic palaeobiogeography (applied to Cenozoic Tethyan molluscs). Palaeogeogr. Palaeoclimat. Palaeoecol. 86: 139-196.
- Prassad, M. 1993. Siwaiik (Middle Miocene) woods from the Kalagarh area in the Himalayan foot hills and their bearing on palaeoclimatic and phytogeography. Palaeobot. Palynol. 1: 49-82.
- Rees, P. M. 1993. Revised interpretations of Mesozoic palaeogeography and volcanic arc evolution in the northern Antartic Peninsula region. Antarct. Sci. 5: 77-84.
- Retallack, G.J. 1992. Middle Miocene fossil plants from Fort Ternan (Kenya) and evolution of African grasslands. Paleobiology 18: 383-400.
- Ruhberg, H. 1985. Die Peripatopsidae (Onychophora). Systematic, Okologie, Chorologie und phylogenetische Aspekte. Zoologica 137: 1-183.
- Sato, S. 1992. On the so-called "Mid-Neogene climatic optimum" in Japan. Fossils 53: 7-19.
- Schuster, R.M. 1983. Phytogeography of Briophyta, p. 463-626. *In* R.M. Schuster (ed.). New Manual of Bryology, vol. I. Hattory Botanical Laboratory, Miyazaki, Japan.
- Schopf, T.J.M. 1980. Paleoceanography. Harvard University, Massachusetts. 341 p.
- Scotese, C.R., L.M. Cahagan & R.L. Larson. 1988. Plate tectonic reconstructions of the Cretaceous and Cenozoic ocean basins. Tectonophysics 155: 27-48.
- Sedgwick, A. 1908. The Distribution and Classification of the Onychophora. Quart.J. Microsc. Sci. 52: 379-406.
- Seyfried, H. & W. Hellmann (eds.). 1994. Geology of an Evolving Island Arc. The Isthmus of Southern Nicaragua, Costa Rica, and Western Panama. Prof. Inst. Geol. Paläont. Univ. Stuttgart 7: 1-433.
- Sieglne-Farkas, A., L. Gyalog & Z. Peregi. 1992. A palynogical study of Middle Eocene formations in the region of Guantanamo and Holguin, Cuba. Magy All Foldt Intez Evi Jel. 1990: 715-735.

- Smellor, M. 1993. Biogeography of Bathonian to Oxfordian (Jurassic) dinoflagellates: Artic, NW Europe and circum-Mediterranean regions. Palaeogeogr. Palaeoclimatol. Palaeoecol. 102: 121-160.
- Srivastava, S.K. & P.L. Binda. 1991. Depositional history of the Early Eocene Shumaysi Formation, Saudi Arabia. Palynology 15: 47-61.
- Tanai, T. 1993. Paleoclimatologic inferences from Tertiary floras. Fossils 54: 35-52.
- Thevenard, F. 1993. The lower Jurassic conifers of the Chaldecoste outcrop, Causses basin (Lozere, France). Palaeobot. Palynol. 78: 145-166.
- Troncoso A., A. 1991. Paleomegaflora of the Navidad formation, Navidad Member (Miocene) in a region of Matanzas (central western Chile). Mus. Nac. Hist. Nat. Bol. 42: 131-168.
- Turner, B.R. 1993. Paleosols in Permo-Triassic continental sediments from Prydz Bay, East Antartica. J. Sediment. Petrol. 63: 694-706.
- Vachon, M. 1954. Répartition actuelle et ancienne des Onychophores ou Péripates. Rev. Gén. Sci. Paris 61: 300-308.
- Varol, O. 1992. Taxonomic revision of the Polycyclolithaceae and its contribution to Cretaceous biostratigraphy. Newsl. Stratigr. 27: 93-127.
- Walker, M.H. 1995. Relatively recent evolution of an unusual pattern of early embryonic development (long germ band?) in a South African onychophoran, Opisthopatus cinctipes Purcell (Onychophora: Peripatopsidae). Zool. J. Linn. Soc. [London] 114: 61-75.
- Walter, H. & S.-W. Breckle. 1986. Ecological Systems of the Geobiosphere. Springer, Berlin (several volumes).
- Xu, Y. L. & S.-Z. Mao. 1992. Cretaceous-Early Tertiary calcareous nannofossils from southern Xizang (Tibet) and their sedimentary environment. Acta Micropalaeontol. Sin. 9: 331-347.
- Yamanoi, T. 1993. Paleotemperature obtained from pollen fossils. Fossils 54: 53-60.